CONTRIBUTION OF REMOTE SENSING TO THE IDENTIFICATION OF GROUNDWATER RESOURCES IN MAYO-DALLAH DEPARTMENT, SOUTH WESTERN CHAD

Abstract: Located in south-western Chad, the Department of Mayo-Dallah has a problem of access to groundwater in terms of quantity due to the complexity of the area (basement). The aim of this study is to extract the lineaments of the Mayo-Dallah Department using remote sensing to identify potential groundwater resources. To achieve this, a satellite image (Radar Sentinel 1-C) was downloaded and subjected to various processing operations. A total of 7193 lineaments were extracted, with lengths ranging from 0.30 to 3Km. 96.7% of the lineaments are small, less than 1Km. Lineaments of 1 to 3Km considered to be major represent 3.29%. The distribution of major lineaments on the directional rosette reveals two main directions E-W and N-S and secondary directions SE-NW, ESE-WNW and NE-SW. The lineament density map shows that the southern and western parts of the study area have good lineament density. These areas have a high water potential and could influence recharge.

Keywords: Chad, groundwater resources, basement, remote sensing, lineaments, Mayo-Dallah.

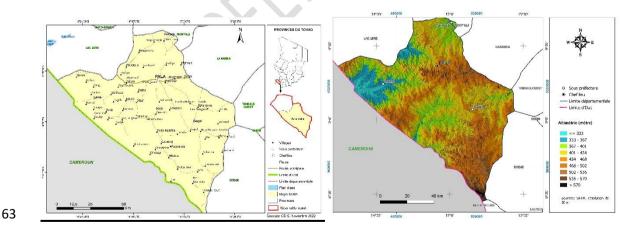
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31	1 - Introduction
32	Groundwater resources are of vital importance to human life. In developing countries,
33	they are the first choice for supplying drinking water to the population, due to their good
34	quality. However, access to these resources is problematic in basement environments. Yet
35	fractured basement aquifers are excellent groundwater reservoirs (Jourda et al., 2006). The
36	Department of Mayo-Dallah, the subject of the present study, is characterized in this context.

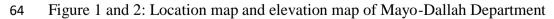
The Department of Mayo-Dallah in south-western Chad belongs to the Sudano-37 Guinean climate zone. It receives an average of 1,000 m of annual rainfall. Despite this 38 rainfall, the area suffers from water problems. Borehole flow rates are often low, and the 39 failure rate is high. It has been observed that in the dry season, the flow of water in boreholes 40 drops considerably and some wells dry up, forcing the population to resort to water of dubious 41 quality. According to several authors, this situation can be explained by the poor choice of 42 drilling sites and also by the lack of hydrogeological knowledge of the area's aquifers. More 43 precise knowledge of fractured aquifers is essential for better location, exploitation and 44 sustainable management of their resources (Lasm, 2000), (Koudou et al., 2014). Remote 45 sensing is considered an ideal tool in the search for water. It is a preliminary method 46 commonly used by hydrogeologists to identify potentially fractured structures considered as 47 underground recharge and flow conduits (Mabee et al 2002), (Liee et Gudmundsson, 2002), 48 (Gleeson et Novakowski, 2009). With this in mind, it is essential to extract lineaments, as 49 their density indicates a high water potential. 50

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2- Presentation of the study area

The Mayo-Dallah Department is located in south-western Chad, in the Mayo-Kebbi 52 Ouest Province, between latitudes 8°57' to 10°15'N and longitudes 15°05' to 15°58'E (figure 53 1). It covers an area of 4,069 km2. It has a Sudano-Guinean climate with two seasons. The dry 54 season runs from November to April, and the rainy season from May to October. Annual 55 rainfall varies from 900 to 1000mm. Average annual temperatures range from 21.8°C to 56 The Department of Mayo-Dallah is home to several temporary watercourses, 57 34.9°C. tributaries of the Mayo-kebbi river, with a sinuous morphology imposed by faults (Doumnang 58 et al., 2021). The topography of the study area is marked by a succession of less accentuated 59 reliefs with numerous small hills. Altitudes range from 300 to 500 m, with an average of 60 380m above sea level. The lowest points are to the west and slightly to the north-west of the 61 study area (figure 1 b). 62





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2.1 - Geological and hydrogeological context

68 The Mayo-Kebbi West Province consists of a Precambrian basement in the north and 69 sedimentary cover formations. The Precambrian basement of Mayo-Kebbi contains three

groups of rocks : the greenstone belts, the mafic and intermediate complex and the granitoid 70 batholith (Mayo-Kebbi batholiths) (Penaye et al., 2005), (Mbaguedjé, 2015). These structures 71 are all oriented in a NNE-SSW direction. They contain ultrabasic rocks (pyroxenites, 72 chloritoschists, talcschists), basic to acidic metaplutonic rocks (gabbros, dioritic gabbros, 73 granodiorites and amphiboloschists) and metavolcanosedimentary rocks (amphibolites, 74 metabasalts, metadolerites and metagrauwackes) affected by greenschist metamorphism [8]. 75 The overlying sedimentary formations correspond to the Lamé series (Cretaceous) and the 76 Continental Terminal (Doumnang, 2006). 77

78 Four phases of deformation mark the evolutionary history of the Mayo-Kebbi basement. The first two phases are E-W shortening and the third and fourth are represented by 79 dextral and senestial detachments respectively (Isseini, 2011). 80

From a hydrogeological point of view, the Mayo-Kebbi basement aquifer is a multi-81 layered aquifer separated by impermeable layers. In the basement, groundwater is located in 82 alterites and fracture zones at an average depth of 40m. The static groundwater level depends 83 more generally on the strength of the alterites. The median specific flow rate is around 0.18 84 m3/h/m. There are also deposits with sandy formations, so permeable and capable of ensuring 85 borehole productivity. Static level depths are moderate, but flow rates are rather low, of the 86 order of 5 m3/h at most, with possible dependence on rainfall. Water levels are generally 10 87 m deep. 88

In sedimentary formations, aquifers are located in the most permeable layers. Water 89 table depths and static levels tend to be moderate, while flow rates are higher and even locally 90 interesting (in excess of 15 m3/h) (Durand, 2003). 91

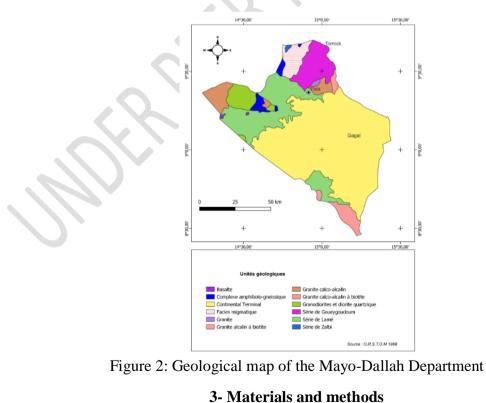
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3.1- Materials

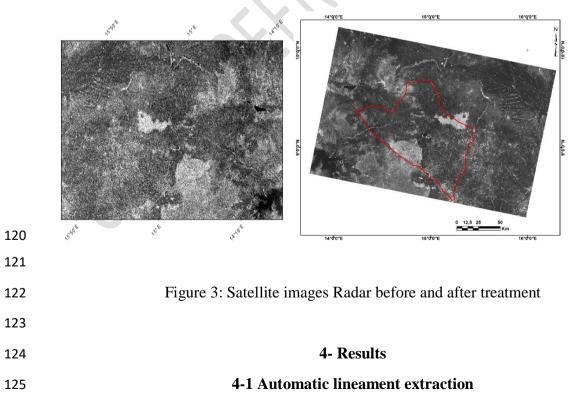
- 97 The material used to achieve our objective is :
- The 1:1,500,000 geological map of Chad,
- Synthetic geological map of the town of Pala,
- Detailed geological map of Pala,
- IGN topographic map at 1/200,000 (sheets NC-33-9, NC-33-10, NC-33-11, NC-33-3 and NC-33-4),
- Satellite imagery (Radar Sentinel 1-C),
- Mayo-Kebbi fracturing map extracted from Landsat 7ETM images, scene p184r053,
 Hydrographic network map extracted from SRTM imagery,
- Borehole data (flow rate and depth).
- 107 These data are processed on several software packages, namely : SNAP, Arc gis, Geomatica108 and Rockwork.

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3.2- Methodology

The methodology adopted for this study consists in applying remote sensing processing 110 methods to extract the lineaments of the study area and to proceed with the control and 111 validation of these lineaments on the one hand, and to show their relationship with 112 113 productivity on the other. To do this, we downloaded an image from the Sentinel 1-C radar sensor satellites taken on March 16, 2024 from the Copernicus Data Space Ecosystem 114 website. The satellite data obtained are dual-polarized (HV + VV). Those taken into account 115 are HV horizontally-vertically polarized. As these are orthorectified images, we directly 116 applied the Lee sigma 7x7 filter. The filtered image is then exported to Geomatica for 117 automatic lineament extraction using the Line algorithm. 118

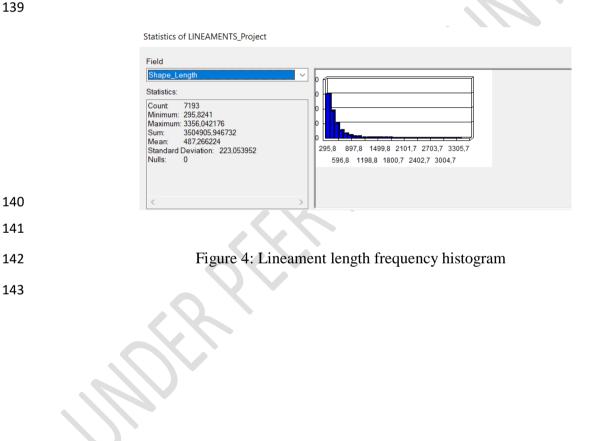
119 The lineaments obtained are then transferred to Rockwork software for directional analysis.



Application of the LINE modulus algorithm extracted 7193 lineaments with sizes 126 ranging from 295m to 3000m. Smaller lineaments (less than 1m) are more numerous (figure 127 4). They account for 96.7% of the total, while lineaments between 1 and 3m account for 128 3.29%. The latter are considered to be the major lineaments. 129

The distribution of lineaments on the directional rosette is fairly homogeneous. The 130 frequencies in number and cumulative length are similar. No family exceeds 10% in number 131 or cumulative length. The N-S and E-W families stand out from the others, with a percentage 132 in number and cumulative length close to 4%. These are the E-W direction (N80° to 100°) 133 and the N-S direction (N170° to 180°) (figure 8). These are followed by N160°-170°, N140°-134 150°, N0°-20°, N70°-80° with almost 3%. The N40-60° and N110-130° directions are poorly 135 represented. We can classify these groups into 5 families: N-S (N0-20°, N160-170), SE-NW 136 (140-150°), ESE-WNW (110-130°), E-W (80-100°), NE-SW (40-60°). 137

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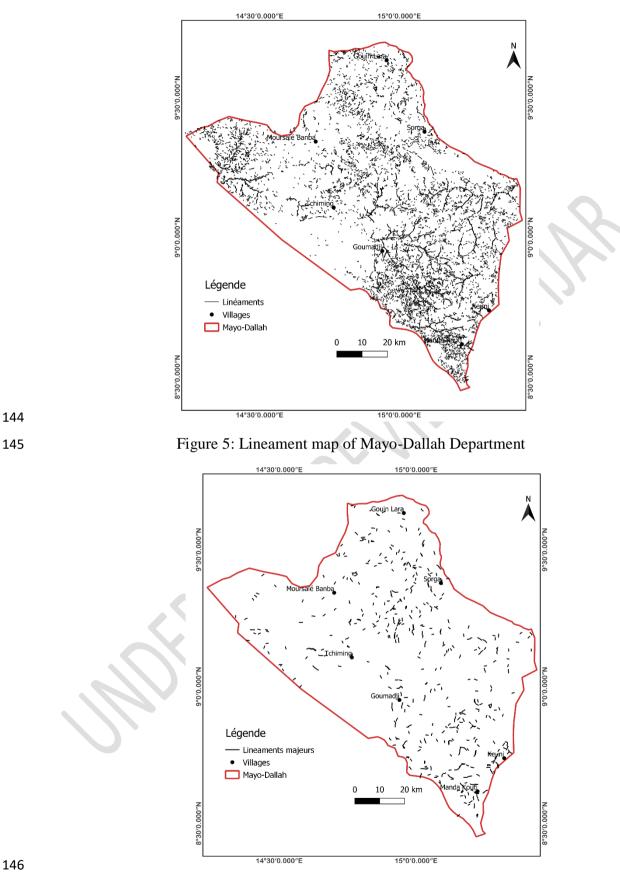
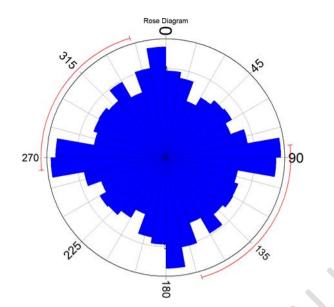
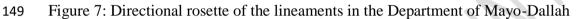




Figure 6: Map of major lineaments in the Department of Mayo-Dallah







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5- Lineament control and validation

Validation enables us to check the reliability of the method used and to assign fracture 152 values to the lineaments obtained. This stage is based on superimposing the major lineaments 153 on the geoscientific data (hydrographic network, fracturing map and high-flow boreholes). By 154 superimposing the lineaments on the hydrographic network of the study area, illustrated in 155 figure 8, we can see that the lineaments follow the watercourses, which seems logical since, 156 according to (Doumnang et al., 2021), the morphology of the watercourses in the Mayo-Kebbi 157 is imposed by the faults. Figure 9 shows a correspondence of more than 50% between the 158 lineaments and the fractures identified, although not all fractures were found in the field due 159 to the poor outcrop in the study area mentioned by previous authors (Penaye et al., 2005), 160 (Isseini, 2011). This result leads to the conclusion that the lineaments in the study area are of 161 tectonic origin. 162

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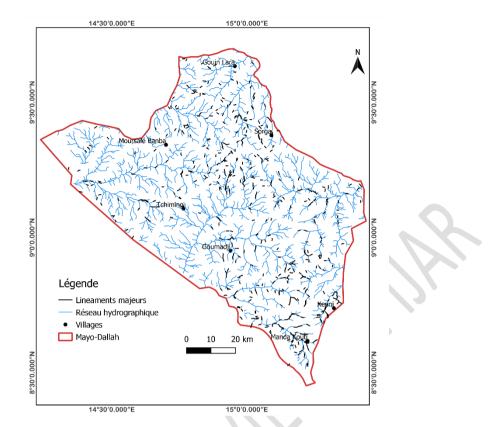
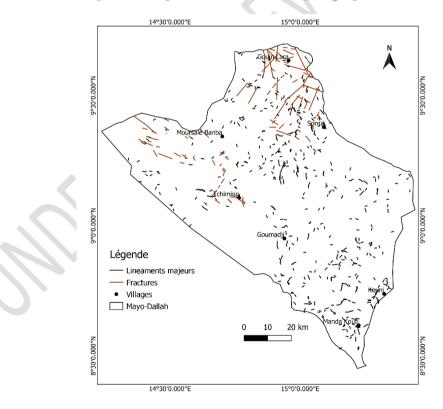




Figure 8: Map of lineaments and hydrographic network



167 Figure 9: Map of major lineaments and fractures in the Department of Mayo-Dallah

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6- Lineament density

171 Density calculates the frequency of lineaments per unit area. The density map shows the

concentration of lineaments per unit area and provides an indication of water potential. Figure
10 shows that the study area is weakly fractured. However, there is good density to the south
and west of the study area.

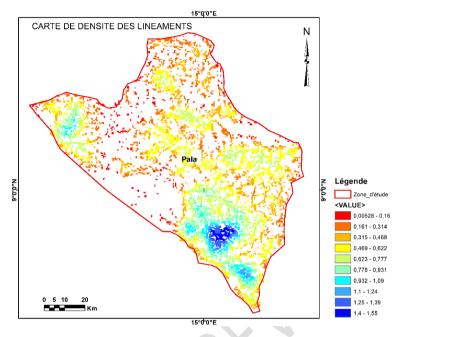




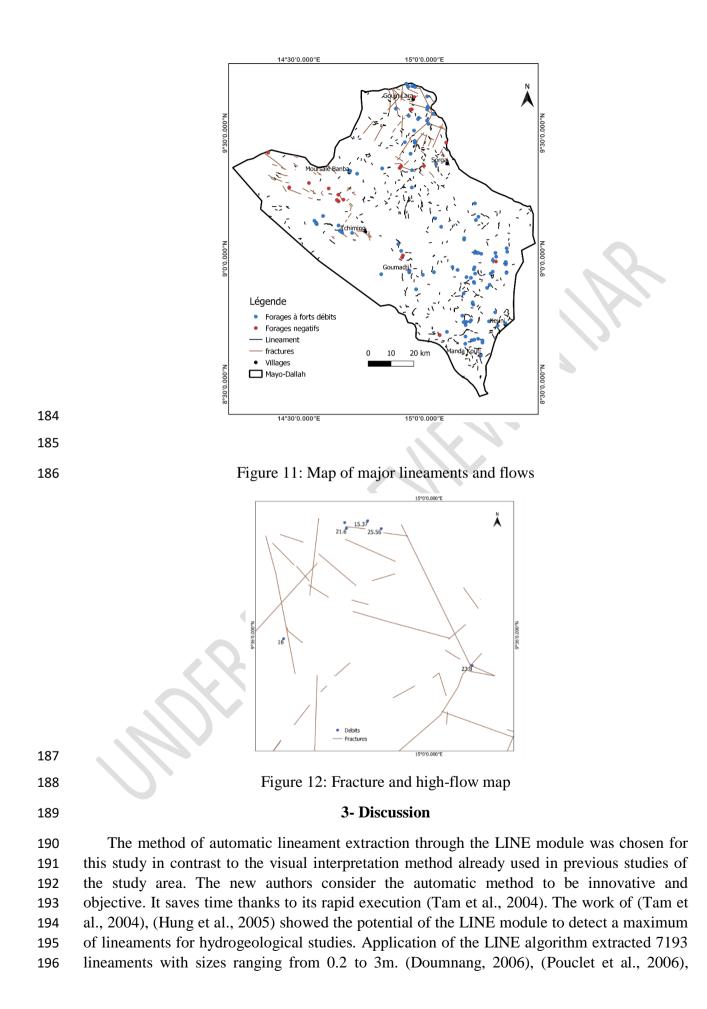
Figure 10: Density map of lineaments

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6- Relationship between lineaments and borehole flow rates

(Krishnamurthy et al., 2000), (Sener, 2005) consider lineaments to be the conduits for groundwater recharge and flow, and their density indicates a high water content. Figure 12 shows the relationship between lineaments and borehole flow rates. Boreholes located near or at the intersection of two or more fractures capture significant flows. Characteristic fractures are highly productive. However, the fractures located to the west of the study area did not yield satisfactory results.

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(Penave et al., 2006) have interpreted the lineament assemblages in the Mayo-Kebbi Province 197 as being related either to the stratification, foliation or schistosity of the rocks, or to their 198 diaclase, fracturation, dyke or vein nature. The distribution of lineaments on the directional 199 rosette shows a homogeneity of lineaments, but the E-W (80-100°) and N-S (170-180°) 200 directions stand out from the others in terms of number and cumulative length. These 201 directions are similar to those of the N170° and E-W (70-100°) normal faults mentioned by 202 (Penaye et al., 2006) in the Mayo-Kebbi basement. These directions are followed by minority 203 directions. A total of five lineament families emerge: E-W (80-100°), N-S (N0-20°, N160-204 170), SE-NW (140-150°), ESE-WNW (110-130°), NE-SW (40-60°). These fracture directions 205 correspond to the directions of the Pan-African (schistosity and shears) and post-Pan-African 206 (Cretaceous fault) structures. The NE-SW direction (N40° and N60°) characterizes the 207 fracture plane of the Mayo-Kebbi greenstone. It is similar to the schistosity directions 208 209 highlighted by (Isseini, 2011; Mbaguedjé, 2015) in the Goueygoudoum series. The SE-NW direction is the same as that of the Doba and Bosso normal faults. This direction is more 210 likely to be "open" and therefore hydrogeologically productive. The ESE-WNW direction is 211 similar to local shear zone directions and structures appear to control certain flow channels 212 (Osinowo et al., 2021). The lineament map of the study area shows all fracture directions in 213 the Mayo-Kebbi basement. However, these lineaments are subject to validation by 214 hydrographic, geological and hydrogeological data from the study area to confirm their nature 215 and existence in the field. The validation of lineaments by the hydrographic network is the 216 very first validation method used by previous authors. The result shows that the lineaments 217 are drawn on the watercourses. This result seems logical since, according to (Vidal et al., 218 2007), the morphology of Mayo-Kebbi streams is imposed by faults. Fractures guide the path 219 220 of rivers and allow lakes to settle in the Mayo-Kebbi Province. Further validation of the lineaments with the geological and fracture maps of Mayo-Kebbi shows that the majority of 221 the lineaments are superimposed on the fractures observed in the field. In view of this result, 222 we can confirm the reliability of the method used and attribute the fracture values to the 223 224 lineaments. The density map shows that the south and west have a good density of lineaments. 225 From a hydrogeological point of view, good fracture densities provide a good understanding of the aquifer. These zones have a high water potential and influence recharge (Mabee, 2002). 226 The coupling of drilling points and fractures shows that there is a relationship between 227 228 productivity and fracture interconnectivity. Boreholes located at fracture intersections or close to kilometre-scale fractures capture significant flows, up to 25m3/s. This hypothesis has been 229 validated by (Biemi et al., 1991), (Savané, 1997), Jourda et al., 2006). However, not all 230 fractures are fed. They may be clogged. This is the case for the boreholes located near the 231 fractures to the west of the study area. In fact, these fractures run in the same direction as the 232 old E-W faults. The old faults are likely to be clogged or vein-filled. The relationship between 233 depth and flow is not significant. It would be important to establish the relationship between 234 alteration depth (thickness) and flow rate to gain more information. 235

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4- Conclusion

The remote sensing tool has made it possible to map the lineaments in the Mayo-Dallah Department and has demonstrated their hydrogeological interest. This study leads to the conclusion that lineaments alone do not guarantee the identification of an aquifer potential. It would be advisable to combine lineaments with other approaches to obtain moredetailed information.

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